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1 **Evaluation of LMS6 and RS41 radiosonde humidity data**  
2 **obtained during YMC-Boreal Summer Monsoon study in**  
3 **2018 and 2020**

4  
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34

## **Abstract**

35

This short article describes humidity data correction based on intercomparison between the two manufacturers' radiosondes with the assessment using precipitable water vapor (PWV) derived from Global Navigation Satellite System (GNSS) signals. In addition, we propose a method to determine whether the same correction procedure can be applied for the case that such intercomparison cannot be conducted.

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During the intensive observation called Years of the Maritime Continent - Boreal Summer Monsoon study in 2018 (YMC-BSM 2018), intercomparison of radiosonde between Lockheed Martin LMS6 and Vaisala RS41-SGP was conducted at Laoag, Ilocos Norte, Philippines from late July to early August 2018. While their mean difference of relative humidity (RH) showed better than 5%, dry bias was confirmed for LMS6 only during clear sky daytime soundings based on the comparison of PWV with that derived from GNSS signals. To use different radiosonde data with the same research-quality, we developed a correction table for LMS6 RH data.

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While a direct intercomparison between different radiosondes and independently developed observational tools such as GNSS-receiver is ideal to evaluate the data quality, it is not always able to be performed. Indeed, we obtained LMS6 radiosonde data at different site at Yap Island, Federated States of Micronesia from another field campaign

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52 YMC-BSM 2020, where any intercomparison could not be conducted. In order to decide  
53 whether the same correction procedure obtained from YMC-BSM 2018 can be applied to  
54 those data, we assessed their similarity based on a relationship between specific humidity  
55 from surface meteorological station data that was obtained independently prior to launch  
56 and radiosonde specific humidity averaged over 300 m from the initial radiosonde  
57 measurement point. This method allowed us to confirm the same behavior between  
58 Laoag data in 2018 and Yap data in 2020, thus we applied our correction method to RH  
59 data in YMC-BSM 2020.

60

61 **Keywords:** YMC; field campaign; radiosonde humidity bias

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## 64 **1. Introduction**

65 Radiosonde is a basic observation tool to measure atmospheric vertical profiles and is  
66 often used in the field campaigns, since it can capture the characteristics of the atmospheric  
67 structure at high vertical resolution (raw data are taken every 4 - 6 m interval). Besides, it is  
68 well-known that additional radiosonde soundings greatly improve analysis fields (Moteki et  
69 al. 2011) and forecast skill (Inoue et al. 2015) through the use of data assimilation. The  
70 international field program, Years of the Maritime Continent or YMC has been carried out  
71 since 2017 (Yoneyama and Zhang 2020). YMC is a unique attempt to understand weather  
72 and climate systems over the Maritime Continent through two different field observational  
73 approaches; intensive observation focusing on specific phenomena with limited observation  
74 period and long-term measurement done by the local meteorological agencies or special  
75 observation tool deployment including mooring system, unmanned autonomous vehicles  
76 and so on. For the most cases, radiosonde is used. The accurate moisture distribution and  
77 variation obtained by radiosonde are crucial to understand dominant atmospheric  
78 convective variabilities such as diurnal cycle of rain, the Madden-Julian Oscillation, and the  
79 boreal summer intraseasonal oscillation over the Maritime Continent and adjacent seas,  
80 where the atmospheric convection is the most active over the world (Neelin and Held 1987;  
81 Sobel et al. 2001; Yamanaka 2016; Yoneyama and Zhang 2020). While the radiosonde data  
82 are believed to be accurate as an in-situ measurement system, it is often reported that their

83 relative humidity data contain erroneous data mainly caused by solar heating of humidity  
84 sensor from the research-quality viewpoint (e.g., Vömel et al. 2007a; Yoneyama et al. 2008).  
85 In addition, their accuracy varies with radiosonde manufacturers (Ciesielski et al. 2010; Nash  
86 et al. 2011). Figure 1 shows a map of radiosonde type used during the two field campaigns  
87 in 2018 and 2020 as shown later. During the campaign, the Vaisala RS41 and Lockheed  
88 Martin LMS6 radiosondes were employed. When we analyze those data, it is important that  
89 any scientific knowledge obtained using those data should be the same regardless of  
90 radiosonde type (LMS6 or RS41). Thus, we need to reduce possible errors as much as  
91 possible. Ciesielski et al. (2012) provided a guideline of general procedure for producing  
92 research-quality radiosonde sounding data from field campaigns. They proposed to adopt a  
93 naming convention for those procedures from level-1 to level-4 in addition to original (raw)  
94 datasets as level-0. The level-1 indicates data in a common format regardless of their  
95 original data types and systems. Then, data processed with automated quality-control  
96 algorithms, which remove apparent errors, are labeled as level-2. The level-3 procedure is  
97 required to identify and reduce systematic errors or biases in the high (original) vertical  
98 resolution data. Level-4 datasets can be obtained from level-3 datasets by producing data  
99 at uniform vertical resolution (e.g., 5 hPa) with quality-control information flags. In these  
100 procedures, they emphasized the importance and difficulty of humidity data bias  
101 identification and reduction. Since the radiosonde data taken as part of the YMC are used

102 by many researchers who did not conduct observations by themselves, data quality should  
103 be evaluated by the researchers who conducted the field experiment. Thus, this is a report  
104 about the quality control of data from those who conducted field campaigns in 2018 and  
105 2020.

106 In this study, we will first show the result of data quality evaluation based on the  
107 intercomparison done during the field campaign at Laoag, Ilocos Norte, Philippines in July -  
108 August 2018 in Section 2. Then, we will apply the same correction procedure developed  
109 from the 2018 campaign onto another field campaign data taken in 2020 at different site,  
110 Yap, Federated States of Micronesia based on the identification of possible bias (Section 3).  
111 This procedure contains a suggestion of humidity data correction for the case that direct  
112 intercomparison cannot be performed during the campaign. A brief summary and remarks  
113 for data quality procedure are provided in Section 4.

114

## 115 **2. Intercomparison during the YMC-BSM 2018**

116 A field experiment to study boreal summer monsoon was conducted as YMC-Boreal  
117 Summer Monsoon study in 2018 (YMC-BSM 2018) at various sites in the tropical western  
118 Pacific region, where unique local phenomena are often observed. The details of the  
119 campaign are found in their website ([http://www.jamstec.go.jp/ymc/campaigns/IOP\\_YMC-  
120 BSM\\_2018.html](http://www.jamstec.go.jp/ymc/campaigns/IOP_YMC-BSM_2018.html)). One key research topic is to study a mechanism of offshore propagation

121 of diurnally-developed atmospheric convections along the coast and its relation to large-  
122 scale atmospheric disturbances known as the northward propagating Boreal Summer  
123 Intraseasonal Oscillation (BSISO). For this purpose, four radiosoundings were performed  
124 per day at Laoag (18.18N, 120.53E), Province of Ilocos Norte, Philippines during July 1 -  
125 August 31, 2018. While the Lockheed Martin LMS6 or Graw DFM-09 radiosonde is usually  
126 used at the Laoag synop airport upper air station, only LMS6 was used during two months  
127 of YMC-BSM 2018 period to avoid inconsistency of radiosonde data. In addition,  
128 intercomparison among different types of radiosonde was conducted to evaluate their  
129 relative differences, so that such information can be used when we use the same type of  
130 radiosonde data taken at different sites. As shown in the next section and in Fig. 1, we also  
131 utilized LMS6 radiosonde data taken at Yap (9.50N, 138.08E) and Palau (7.37N, 134.54E),  
132 as well as Vaisala RS41-SGP (RS41) onboard the research vessel (R/V) Mirai (12.00N,  
133 135.00E) and Legazpi (13.15N, 123.73E), Province of Albay, Philippines, during the YMC-  
134 BSM 2020 campaign. In this study, we will focus on the evaluation of LMS6 radiosonde  
135 humidity data, which are used for both the campaigns. Since the processing software of  
136 LMS6 is different for the various sites and it may also affect the results, their software version  
137 information is provided in Table 1. As shown, while LMS6 transmitter is used at Yap, they  
138 use the U.S. National Oceanic and Atmospheric Administration (NOAA)'s original software  
139 "Radiosonde Replacement System (RRS) Workstation Software (RWS)" (NOAA, 2018). In

140 addition, we used relatively new sensors (within a year from their production) to avoid  
141 possible errors from sensor aging.

142 To evaluate LMS6 radiosonde humidity data, intercomparison between LMS6 and RS41  
143 was performed by launching both radiosondes simultaneously for 20 times during July 27 -  
144 August 3, 2018 (Table 2). It should be noted that while we attached them 1m apart using a  
145 plastic rod and launched using a single balloon at 0300 and 0900 UTC, we launched them  
146 simultaneously but separately using two balloons at 0000 and 0600 UTC. This is because  
147 we kept the same launching procedure for 6-hourly sounding using LMS6 during the entire  
148 2-month YMC-BSM 2018 campaign period. Unfortunately, two soundings had failed to  
149 obtain data correctly for intercomparison, thus 18 sounding data (13 in daytime, 5 in evening)  
150 were used. In addition, during this period a Global Navigation Satellite System (GNSS)  
151 receiver was installed at the Laoag weather station to measure precipitable water vapor  
152 (PWV; Rocken et al. 1993). The procedure to derive PWV from GNSS signal is the same as  
153 that shown in Fujita et al. (2008, 2020), who reported the mean difference with radiosonde-  
154 derived PWV is better than 1mm with a root mean square error of 2-3 mm.

155 The result of intercomparison between LMS6 and RS41 is shown in Fig. 2. It should be  
156 noted that the intercomparison was performed under clear sky condition as shown in Table  
157 2. Note that here we use the term “clear sky” not as cloud amount “0” but as within fewer  
158 clouds condition ranging 1-3 except one case of 4 at sounding number 15 in Table 2. Relative

159 humidity (RH) profiles of LMS6 show dry bias comparing to that of RS41 for entire  
160 troposphere (mean difference between the two radiosondes in the 1000-200 hPa layer is  
161 3.9%, and the difference increases with height in the upper troposphere above 200 hPa  
162 level) during daytime sounding (Fig. 2a), while such difference cannot be found in the  
163 evening soundings (Fig. 2b). Actually, we cannot conclude that LMS6 has dry bias only from  
164 this figure, as RS41 may have a wet bias. Note that if we assume that RS41 data is correct  
165 (no wet nor dry bias), the difference of 3.9% is ranged within the LMS6 specified accuracy  
166 ( $\pm 5\%$ ), suggesting that such a difference is acceptable from the operational-use viewpoint.

167 Having said that, the same quality among different types of radiosonde is desirable  
168 especially when the researchers use them to calculate a moisture flux in a certain domain  
169 or discuss humidity variation using the different radiosonde data. As one index which can be  
170 obtained independently from the radiosonde measurement, a comparison of PWV derived  
171 from two sources; radiosonde and GNSS, is useful. Indeed, previous studies demonstrated  
172 the usefulness of GNSS-derived PWV for assessing radiosonde humidity data validation  
173 (Wang and Zhang 2008; Wang et al. 2013). Figure 3 shows PWV differences between  
174 GNSS-PWV and Radiosonde-PWV for LMS6 (Fig. 3a) and RS41 (Fig. 3b), respectively. We  
175 can confirm that the difference between GNSS and RS41-derived PWV is smaller than 1  
176 mm (e.g., mean difference at 14 LST is -0.92 mm with a standard deviation of 1.1 mm)  
177 regardless of the time of the day, while LMS6 shows large dry bias in daytime (the difference

178 at 14 LST is -3.8 mm with a standard deviation of 1.2 mm). The difference found in RS41  
179 and GNSS is comparable to the previous works (e.g., Realini et al. 2014), which compared  
180 GNSS-PWV with Vaisala RS92-derived PWV, but the data consistency between RS92 and  
181 RS41 are confirmed by several works (e.g., Jensen et al. 2016; Kawai et al. 2017). Based  
182 on this result, we decided that RS41 humidity data is more reliable than that obtained by  
183 LMS6, and we will use RS41 as a reference for RH data correction of LMS6.

184 Here, before describing correction procedure, we discuss possible causes of the biases  
185 briefly. The clear contrast between the daytime and evening suggests that the differences  
186 between the two radiosonde types in the daytime soundings are caused by daytime solar  
187 radiation (Vömel et al. 2007a), although LMS6 has a sunshade to protect from solar radiation  
188 as much as possible. Since the difference can be found from the ground surface, it is  
189 possible to speculate that exposure under the sunshine prior to launch about 5 minutes or  
190 more might result in some portions of this dryness. Besides, it is also possible to speculate  
191 that sunshade could not fully prevent from solar radiation during ascent with slight inclination  
192 rather than upright. Larger differences above 200 hPa might reflect the fact that all  
193 radiosondes launched during the intercomparison period encountered strong winds over 10  
194  $\text{m s}^{-1}$ (not shown), which might result in such sunshade could not prevent from the solar  
195 radiation due to inclination. There are also several other possible reasons to explain the  
196 difference. For example, we cannot exclude a possible error source by the slow sensor

197 response at low temperatures, known as a time-lag error (Miloshevich et al. 2004). However,  
198 it is impossible to quantify the biases due to the time-lag error, as we do not have information  
199 about the sensor time constant which is determined by the manufacturer on the basis of  
200 dedicated measurements in a laboratory environment. Instead, what we can do is to  
201 minimize their gap between RS41 and LMS6 in producing seamless humidity datasets  
202 based on the fact that the differences exist only in clear sky daytime soundings.

203 As for correction schemes, several methods have been proposed and used for previous  
204 field campaigns (e.g., Nuret et al. 2008; Ciesielski et al. 2014). If we could obtain the  
205 simultaneous sounding data using chilled-mirror type hygrometers such as Snow-white  
206 (Fujiwara et al. 2003) or cryogenic frost-point hygrometer (Vömel et al. 2007b) with enough  
207 sample numbers, a polynomial fitting with those data profiles might be better. However, we  
208 did not launch such hygrometers during the 2018 campaign. Thus, by assuming that RH of  
209 RS41 is better than that of LMS6 during the 2018 campaign as mentioned above, we fitted  
210 LMS6 RH data to that of RS41 using the cumulative distribution function (CDF) matching  
211 method, which produces a look-up table for humidity bias correction (Nuret et al. 2008;  
212 Ciesielski et al. 2009). Since we could not find significant difference between 1100 LT and  
213 1400 LT sounding cases, we simply calculated CDF matching using all 13 samples for  
214 daytime sounding. RH bias as a function of RH and temperature is shown in Fig. 4. Then,  
215 bias correction is performed onto the RHs and its result is shown in Fig. 5, which

216 demonstrates that the humidity correction performed well for entire troposphere. Figure 6,  
217 which shows the frequency distribution of occurrence of RH for original/corrected values,  
218 provides the same information with Fig. 5, but it also tells how this correction works. The  
219 peak of frequency in the higher RH range, which is located at 65-75% in the original LMS6  
220 data, shifts to 75-85% in the corrected one, while another peak around 20-30% in the original  
221 LMS6 disappears in the corrected distribution. The former may capture moist condition in  
222 the lower troposphere over the warm pool, while the latter corrected dry bias in the upper  
223 troposphere mainly above 200 hPa shown in Fig. 2a.

224

### 225 **3. Application of correction scheme to another field campaign data**

226 Although we obtained the quality-controlled LMS6 humidity data during the YMC-BSM  
227 2018 campaign period, it does not guarantee such correction is valid for other LMS6 data,  
228 as the system has been improved by the manufacturer's effort and minor changes of the  
229 system is often done without notice to the users.

230 In the boreal summer in 2020 (August - early September), another YMC field campaign  
231 called YMC-Boreal Summer Monsoon study in 2020 (BSM 2020) was conducted. During  
232 this campaign period, a radiosonde network was formed by three land-based sites and the  
233 R/V Mirai (indicated in blue in Fig. 1). While RS41 was used at Legazpi and onboard the  
234 R/V Mirai, LMS6 was used at two sites; Yap, Federated States of Micronesia, and Palau. At

235 these land-based sites, the twice daily operational radiosoundings were increased four times  
236 per day a six-hour interval during the campaign period. As for LMS6 data at these sites, is it  
237 possible to use their data without any correction? Or, is it acceptable to adopt the same  
238 correction procedure obtained from the YMC-BSM 2018 Laoag site as shown in the previous  
239 section? All what we obtained from those sites are only original high-resolution (every 1  
240 second) radiosonde data, and we did not conduct any intercomparison there.

241 One idea to answer to above is the use of surface meteorology data taken independently  
242 using their surface weather station as initial ground surface data prior to launch. It is  
243 expected that specific humidity at the surface measured by the weather station and in a well-  
244 mixed boundary layer measured by the radiosonde should be close each other. For example,  
245 Fitzjarrald and Garstang (1981) suggested the difference are between 1.0 - 1.5 g/kg from  
246 tropical Atlantic Ocean experiment. However, since the LMS6 data obtained at Laoag during  
247 the YMC-BSM 2018 showed dry bias even in the boundary layer (Fig.2a), it is expected that  
248 their difference may be larger. Figure 7 shows scatter plot of the specific humidity measured  
249 at the surface ( $Q_{SFC}$ ) and that averaged over 300m from the radiosonde initial data point  
250 ( $Q_{BL}$ ). Mean difference between  $Q_{SFC}$  and original  $Q_{BL}$  for daytime sounding (2.31  
251 g/kg) is higher than that for evening sounding (0.88 g/kg), and the contrast (1.43 g/kg) is  
252 comparable to the LMS6's dry bias of about 5% with respect to RS41 in the lower  
253 troposphere. The RH correction for daytime soundings improves the mean difference

254 between  $Q_{SFC}$  and  $Q_{BL}$  to 1.14 g/kg that is close to that for evening soundings.

255 Then, we also calculated the same scatter plot for all data obtained during YMC-BSM  
256 2018 period (Fig. 8). Here, it should be noted that although similar comparison between  
257 surface and boundary layer was used to detect dry bias and correction was applied to all  
258 data for the TOGA-COARE (Tropical Ocean Global Atmosphere - Coupled Ocean-  
259 Atmosphere Response Experiment) field experiment (Zipser and Johnson 1998; Lucas and  
260 Zipser 2000), it is different from our case. For TOGA-COARE data, solar radiation-induced  
261 error is one of error sources, but the chemical contamination between the humidity sensor  
262 and the radiosonde package material mainly caused dry bias for Vaisala RS80 radiosonde  
263 data (Wang et al. 2002). Namely, such dry bias appeared in all sounding data. On the other  
264 hand, the dry bias studied here is mainly caused by daytime solar radiation. Thus, a  
265 correction scheme developed here can be applied only to data taken during daytime. In  
266 addition, when the daytime data are obtained under heavy clouds or rainy days, it may not  
267 be necessary to be corrected. As noted in Table 2, the intercomparison was conducted only  
268 under clear sky condition. In other words, we should not apply the correction scheme  
269 obtained from the intercomparison at YMC-BSM 2018 to the data obtained in cloudy/rainy  
270 days. In this study, we identified cloudy/rainy days using a threshold of 93% RH following to  
271 Zuidema (1998). When the data show RH higher than 93% and have enough thickness  
272 (thicker than 30 hPa) between 800 hPa and 300 hPa, we identified the sounding site is

273 covered with cloud, whose threshold thickness (30 hPa) was determined based on several  
274 trials to detect the most similar value with that of non-daytime period. Note that, however,  
275 their difference is not changed much from those layer-thickness selections, because it is  
276 often covered with thick layers due to rain when several points show the values higher than  
277 93% RH (not shown). We confirmed that all daytime soundings shown in Table 2 were not  
278 categorized in this cloudy/rainy condition defined using RH threshold. Also, note that the  
279 range starting from 800 hPa was chosen to surely avoid the top of boundary mixed layer,  
280 where high RH often appears, and cloudy layer from spotted/scattered shallow clouds. Thus,  
281 in Fig. 8, we plotted specific humidity at the surface and 300 m-layer mean before/after  
282 correction for three different cases: clear sky daytime, cloudy/rainy daytime, and others  
283 (evening - night). Those are indicated in different colors and shapes. Namely, only data taken  
284 during clear sky daytime were corrected (indicated in red and magenta). Indeed, we  
285 confirmed that the mean difference between  $Q_{SFC}$  and  $Q_{BL}$  during cloudy/rainy daytime  
286 (1.30 g/kg) is significantly lower than that of clear sky daytime (2.20 g/kg), and rather very  
287 close to that of the others (1.23 g/kg).

288 Then, the similar plot was calculated for the data obtained at Yap during August 6 -  
289 September 9, 2020 (Fig. 9). We can confirm the similar difference and relationship for Yap  
290 LMS6 data. The mean difference between  $Q_{SFC}$  and  $Q_{BL}$  for clear sky daytime  
291 soundings (2.49 g/kg), which is close to that for Laoag BSM2018 case (2.20 g/kg), is

292 significantly higher than that for cloudy/rainy daytime soundings (1.32 g/kg), which is also  
293 close to that for the Laoag case (1.30 g/kg). Namely, it is possible to recognize that LMS6  
294 used at Yap may have similar dry bias with that in YMC-BSM 2018 and the same correction  
295 procedure can be applied, despite the data processing software is different (Table 1).  
296 According to the data policy of the YMC, we are required to provide quality-controlled data  
297 to the public within one year from the completion of the field campaign. Thus, when we  
298 release those YMC-BSM 2020 radiosonde data, they will be labeled as levels-3 and 4  
299 (Ciesielski et al. 2012) with notation of such humidity correction based on YMC-BSM 2018  
300 intercomparison result.

301

#### 302 **4. Summary and concluding remarks**

303 This article reported the evaluation of LMS6 radiosonde humidity data obtained during the  
304 YMC-BSM 2018 campaign at Laoag, Ilocos Norte, Philippines based on the intercomparison  
305 with simultaneous observation with RS41 radiosonde and comparison of PWV derived from  
306 radiosonde and GNSS signal. LMS6 showed a dry bias of approximately 4%RH or so  
307 comparing to RS41 in the entire troposphere with higher difference above 200 hPa level.  
308 Such differences were confirmed only during the clear sky daytime soundings and  
309 recognized that they were caused by humidity sensor heating due to solar radiation. Since  
310 the comparison of PWV derived from radiosondes and GNSS showed that RS41-derived

311 PWV was closer to GNSS-derived one, we decided to correct LMS6 bias using RS41 data.  
312 The CDF matching method was applied to deduce correction values of LMS6 RH data  
313 against RS41 as a function of observed RH and temperature.

314 In addition to the YMC-BSM 2018 data, we obtained LMS6 data from another field  
315 campaign YMC-BSM 2020. In this study, we showed the case of data taken at Yap,  
316 Federated States of Micronesia during August - September 2020. Note that although we  
317 obtained LMS6 data from Palau, currently we have not obtained their surface meteorological  
318 data taken independently. Thus, we focused only Yap data in this study. However, as noted  
319 in Table 1, since Palau adopts the similar system with Laoag, we expect the current study  
320 can be applied to them.

321 Since we do not have any intercomparison between LMS6 and RS41 data during the  
322 YMC-BSM 2020 campaign period, we proposed to use a relationship between the surface  
323 specific humidity data and mean specific humidity data averaged over 300m in the boundary  
324 layer as one possible index to judge whether the correction scheme developed for YMC-  
325 BSM 2018 can be used for the 2020 case or not, since the surface value is obtained  
326 independently from the radiosonde at most stations. Based on the qualitative similarity (their  
327 mean difference is the same order for both campaigns, and at least we confirmed that the  
328 mean difference between  $Q_{SFC}$  and  $Q_{BL}$  for clear sky daytime soundings was  
329 significantly higher than that for cloudy/rainy daytime soundings as Laoag case showed),

330 we decided to apply the same correction scheme onto the Yap LMS6 sounding data. This  
331 suggests that it is possible to apply the same procedure to other sites, where  
332 intercomparison or any other evaluation mechanism have not been established, once the  
333 correction values for the same type of radiosonde have been determined at other sites.  
334 Besides, we used RH values to detect cloudy/rainy conditions during the sounding (judged  
335 as cloudy/rainy, if 93% or higher RH appears over 30 hPa thickness between 800 hPa and  
336 300 hPa), so that the humidity correction associated with solar radiation in daytime can be  
337 applied only to affected data under clear sky days.

338 The datasets used in this study are available from the YMC data archive site at  
339 [http://www.jamstec.go.jp/ymc/ymc\\_data.html](http://www.jamstec.go.jp/ymc/ymc_data.html), and a look-up table for LMS6 is also available  
340 in the same site and Supplement. It should be noted again that this look-up table for LMS6  
341 can be used only for the data taken during clear sky daytime.

342 Currently, the authors have also been conducting to develop correction scheme for other  
343 radiosonde types such as Graw obtained at Laoag. We believe those efforts should be done  
344 in collaboration with the local meteorological agencies, so that not only the field campaign  
345 data but also operational routine data can be used for weather and climate studies with high  
346 research-quality.

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492

Site	Transmitter Type	Ground System Software Version
Laoag	Vaisala RS41-SGP	MW41 2.3.0
	Lockheed Martin LMS6	LMG6 Win9000 6.3.1
Yap	Lockheed Martin LMS6	NOAA RWS 3.4.0.2
Palau	Lockheed Martin LMS6	LMG6 Win9000 6.6.0

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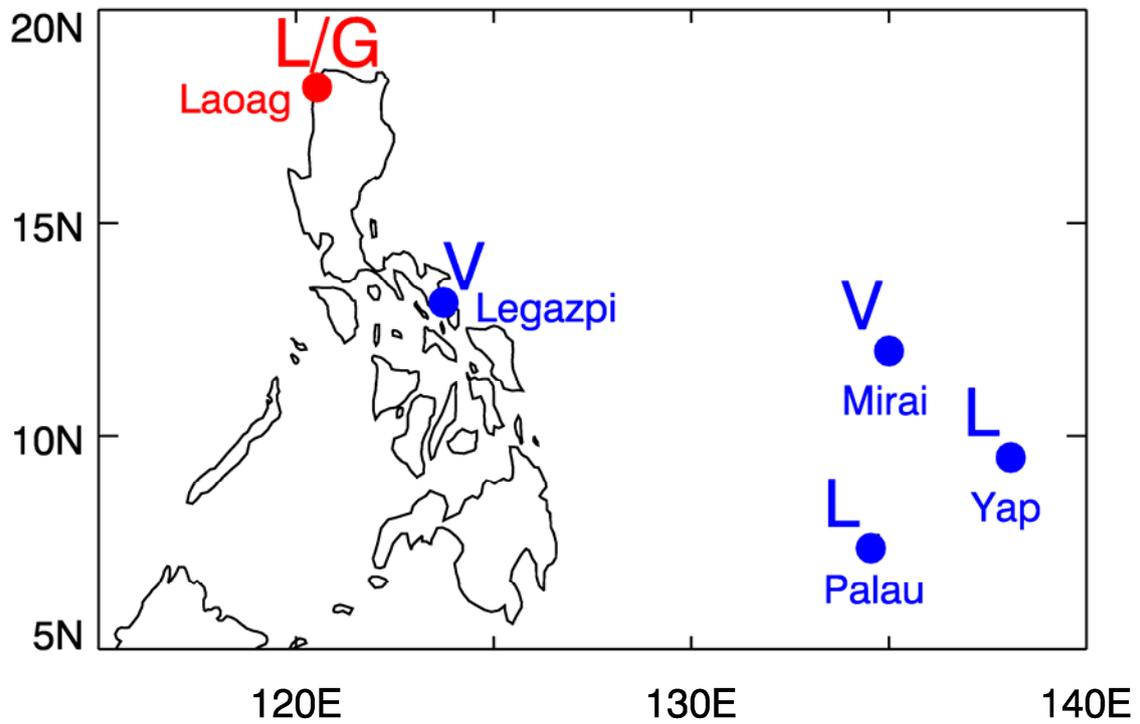
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 497 PAGASA Laoag station, and they are expressed in okta (eighths of the sky).  
 498

No.	Nominal Time (UTC)	Actual launch time (LT)	Remarks
1	2018.07.27 09:00	2018.07.27 16:36	Cloud amount (3)
2	2018.07.28 03:00	2018.07.28 10:33	Cloud amount (1)
3	2018.07.28 09:00	2018.07.28 16:32	Cloud amount (3)
4	2018.07.29 03:00	2018.07.29 10:35	Cloud amount (3)
5	2018.07.29 06:00	2018.07.29 13:34	Cloud amount (3)
6	2018.07.29 09:00	2018.07.29 16:35	Cloud amount (1)
7	2018.07.30 03:00	2018.07.30 10:44	Cloud amount (2)
8	2018.07.30 06:00	2018.07.30 13:31	Cloud amount (1)
9	2018.07.30 09:00	2018.07.30 16:35	Cloud amount (1)
10	2018.07.31 03:00	2018.07.31 10:35	Cloud amount (2)
11	2018.07.31 06:00	2018.07.31 13:33	Cloud amount (1)
12	2018.07.31 09:00	2018.07.31 16:31	Cloud amount (2)
13	2018.07.31 12:00	2018.07.31 19:31	Cloud amount (2) Data are unavailable
14	2018.08.01 03:00	2018.08.01 10:40	Cloud amount (3)
15	2018.08.01 06:00	2018.08.01 13:30	Cloud amount (4)
16	2018.08.02 00:00	2018.08.02 07:38	Cloud amount (1) Data are unavailable
17	2018.08.02 03:00	2018.08.02 10:32	Cloud amount (2)
18	2018.08.02 06:00	2018.08.03 13:30	Cloud amount (1)
19	2018.08.03 03:00	2018.08.03 10:31	Cloud amount (1)
20	2018.08.03 06:00	2018.08.03 13:30	Cloud amount (3)

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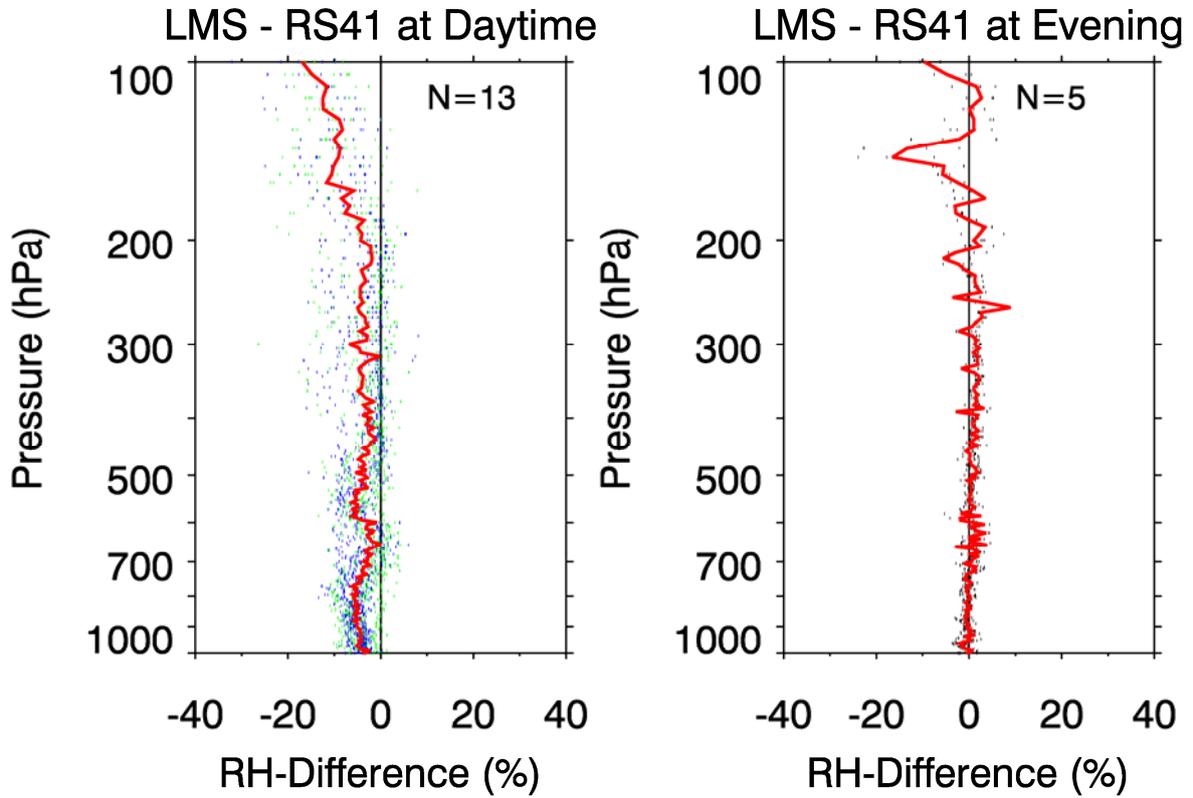


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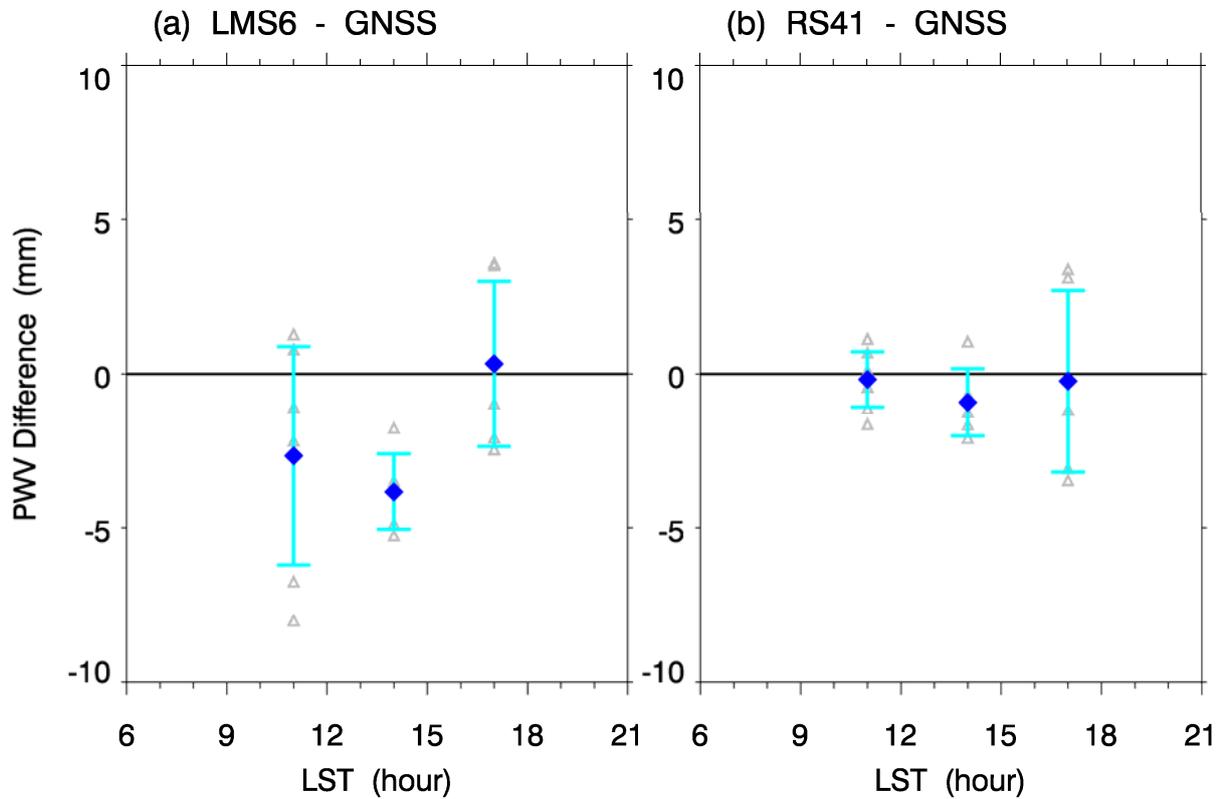
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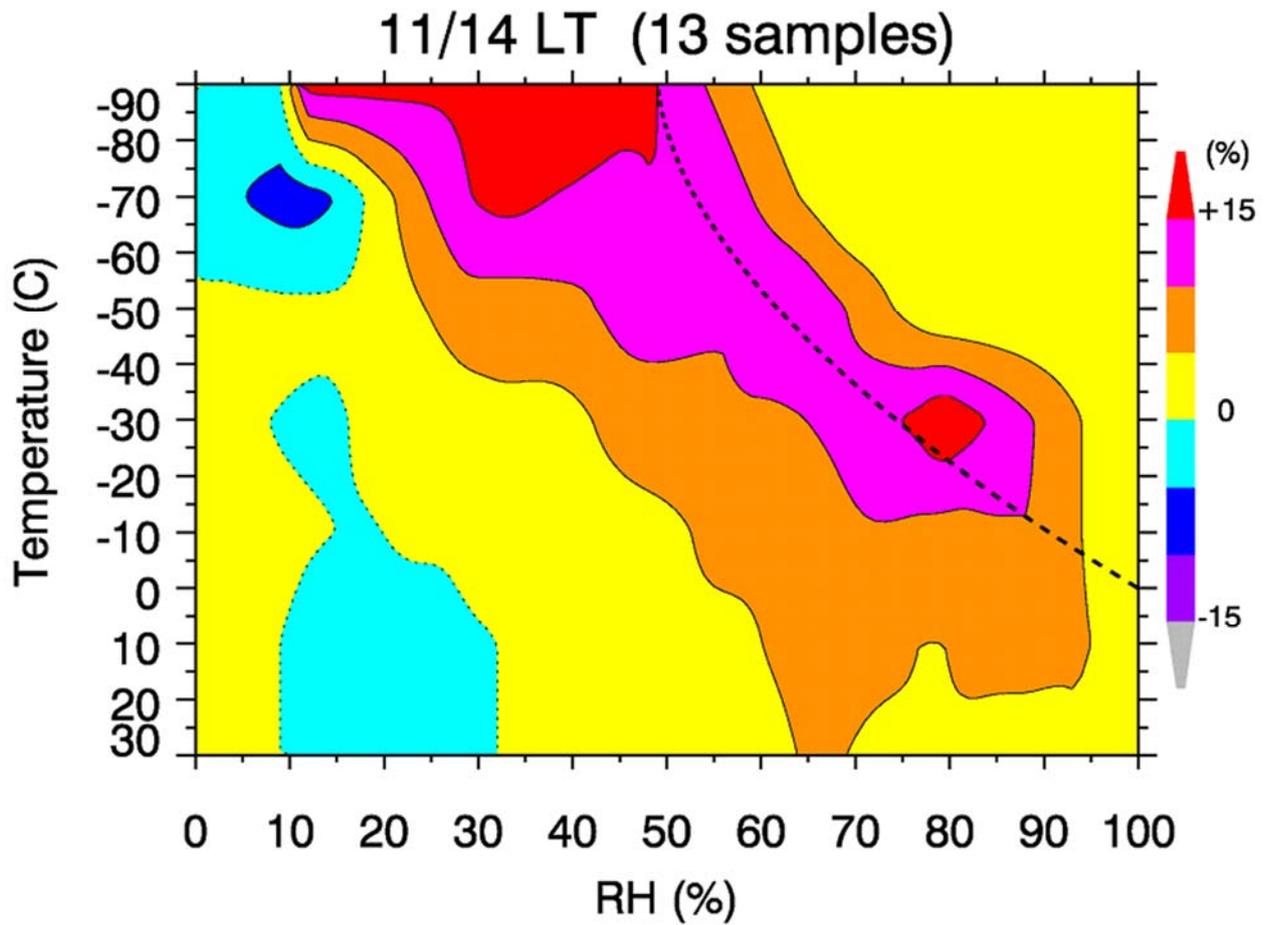
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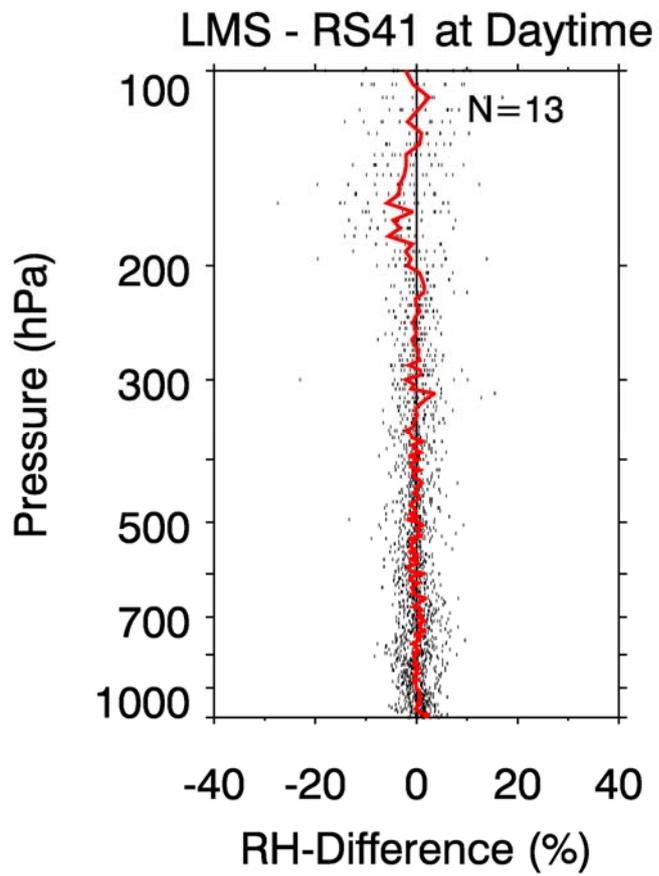
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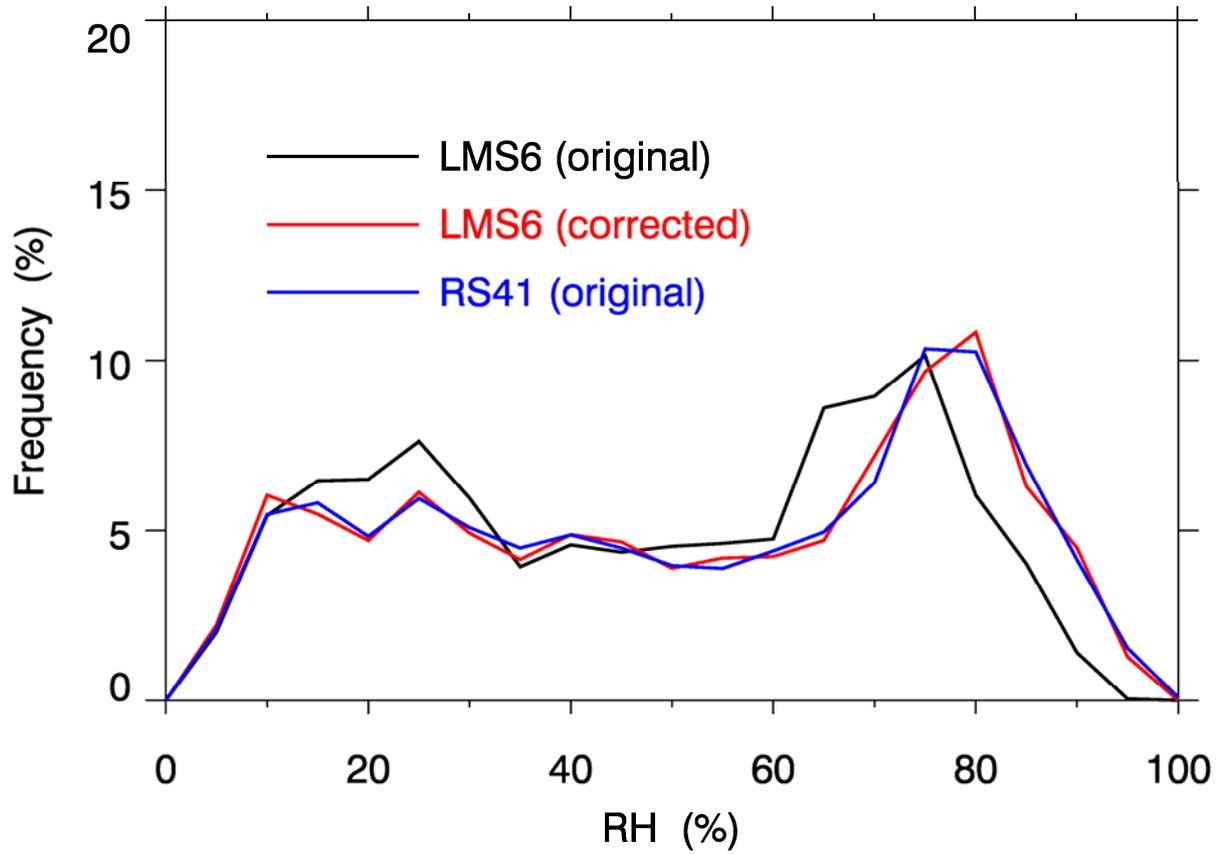
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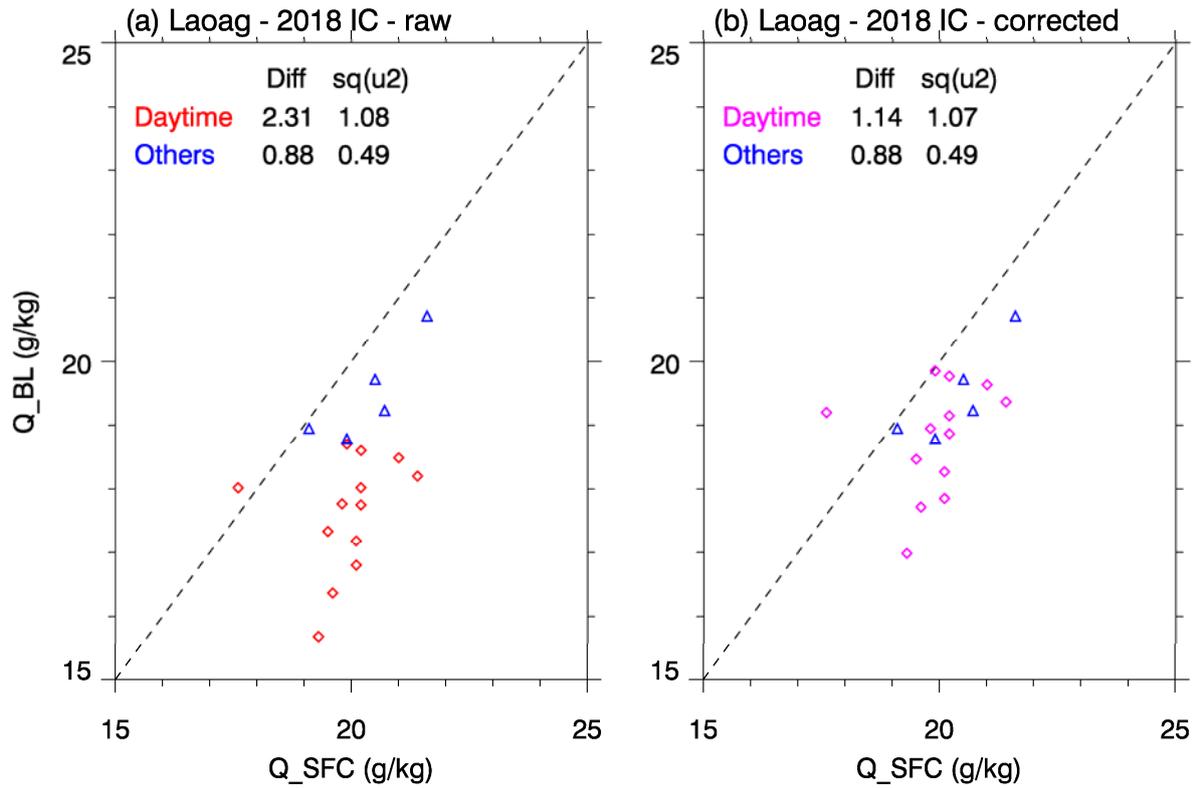
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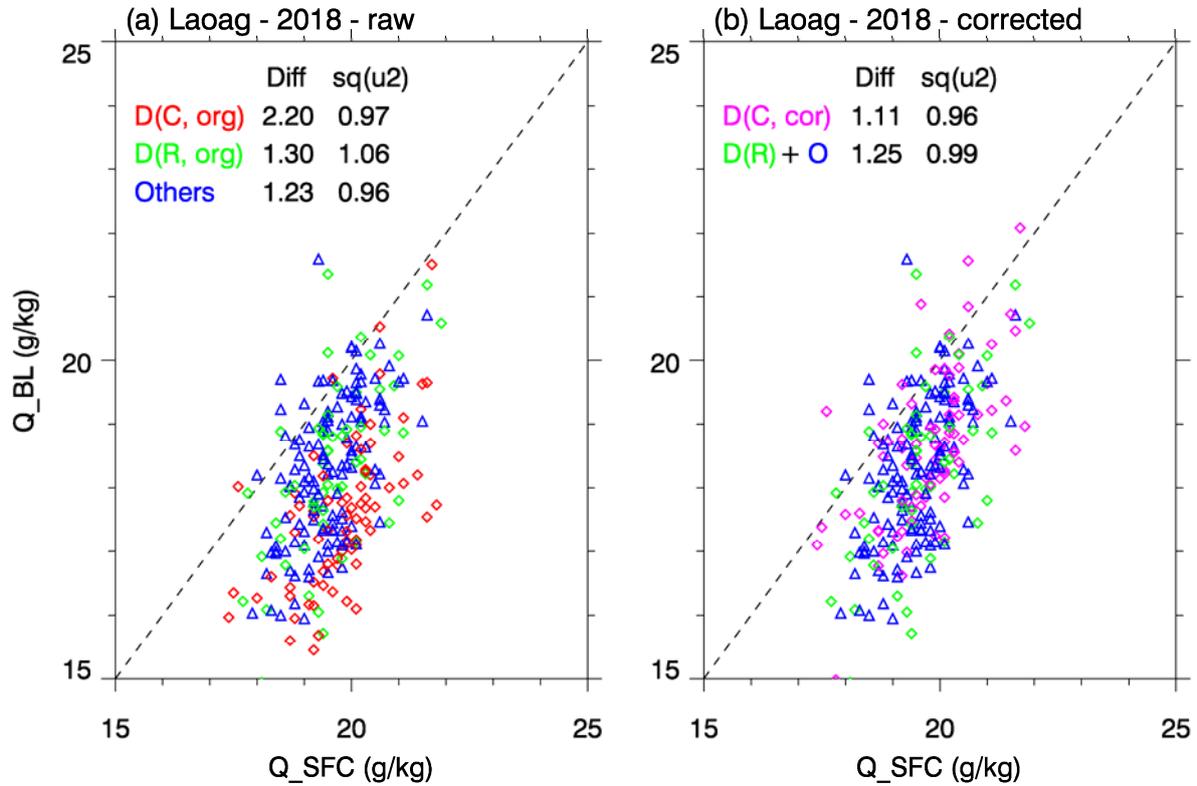


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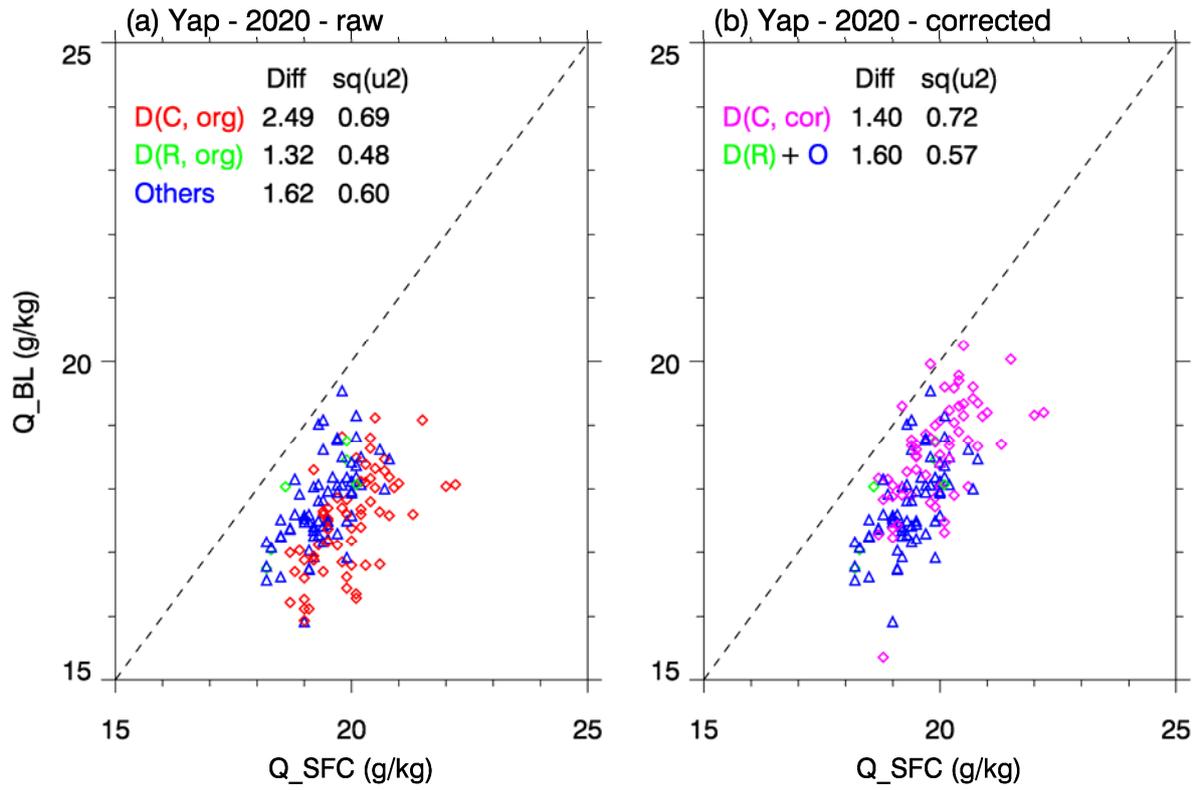
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